

New radiolarian biostratigraphic age constraints on Middle Triassic basalts and radiolarites from the Inner Hellenides (Northern Pindos and Othris Mountains, Northern Greece) and their implications for the geodynamic evolution of the early Mesozoic Neotethys

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Abstract The Avdella Mélange in the northern Pindos Mountains and its equivalent formation, the Loggitsion Unit in the Othris Mountains expose early Mesozoic (Mid-Late Triassic) oceanic fragments beneath the Western Greek Ophiolite Belt of the Inner Hellenides, Northern Greece. The mélange consists of locally interfingering blocks and slices of ribbon radiolarite, radiolarian chert and pillow basalt and is usually overthrust by Jurassic ophiolites. New Middle and Upper Triassic radiolarian biostratigraphic data are presented from radiolarites and basalt-radiolarite sequences within mélange blocks. Pillow basalts associated with the radiolarites provide clues to the opening of the Neotethyan ocean basin. The radiolarians indicate a Middle Triassic age (latest Anisian, probably early Illyrian), which is

documented for the first time in the northern Pindos Mountains. The new radiolarian biostratigraphic data suggest that rift-type basalt volcanism already began in pre-Ladinian time (late Scythian?—Anisian). These basalts were then overlain by Upper Anisian to Carnian (?Norian) radiolarites.

Keywords Triassic · Hellenides · Greece · Radiolarians · Basalts · Radiolarites

Introduction

Early rift-related submarine volcanics and related deep-sea sedimentary units, including radiolarites of the Western Tethys orogenic belt, provide important clues to the geodynamic evolution of the early Mesozoic Neotethys. Dating of intercalated radiolarites and basalts can help to constrain the initial opening of Neotethys. Most geodynamic reconstructions suggest onset of ocean seafloor spreading during Middle–Late Triassic time (e.g. Jones and Robertson 1991; Jones et al. 1992; Danelian et al. 2000; Danelian and Robertson 2001; Bortolotti et al. 2001, 2006; Robertson 2007). Previously, the oldest biostratigraphic ages (upper Illyrian, Anisian) came from similar mélanges that extend from the Inner Dinarides to the Inner Hellenides. An upper Illyrian (Anisian) age was specifically reported from the Mirdita Zone in Albania (Chiari et al. 1994, 1996; Kellici et al. 1994; Marcucci et al. 1994; Bortolotti et al. 2006; Gawlick et al. 2008), where radiolarian cherts are associated with MOR-type basalts and also from basalt-radiolarian chert sequences in the Othris Mountains (Bortolotti et al. 2008).

The aim of this study is to present new radiolarian biostratigraphical age data from several localities in the northern Pindos Mountains, where radiolarites are intercalated with and directly overlie basalts. In addition, we

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give results from one locality in the Othris Mountains, where the contact relations are less well defined but where the oldest known radiolarian fauna also occur. We also compare our new data with similar Neotethyan units in southeast Slovakia and northeast Hungary and more widely throughout the Inner Dinarides and Inner Hellenides as a whole. The results allow an improved reconstruction of the early evolution of the Mesozoic Tethys.

Geological setting

The Pindos and Othris Mountains, where the radiolarian data come from, form parts of the Western Greek Ophiolite Belt (Fig. 1). Both areas consist of thrust sheets and *mélange* (Jones and Robertson 1991) and are assigned to the Maliak Zone (Ferrière 1976, 1982) or to the Pindos-Othris zone (Jones and Robertson 1991; Rassios and Smith 2000). The regional tectono-stratigraphy comprises the following units: 1. Triassic rift-related basalts and sediments within tectonic-sedimentary *mélange* (Avdella *Mélange* in the northern Pindos Mountains: Jones and Robertson 1991; Kemp and McCaig 1984; Pe-Piper 1998; Saccani et al. 2003; Loggitsion Unit in Othris Mountains: Ferrière 1982), 2. Middle Jurassic MORB-type and ophiolitic sequences (Dramala Complex in the northern Pindos Mountains: e. g. Capedri et al. 1980; Jones and Robertson 1991; Metallion, Fourka and also some other units in the Othris Mountains: Ferrière 1982), 3. Mesozoic pelagic, platform-related sedimentary sequences, and turbiditic slope sediments (Dio Dendra Group: in the northern Pindos Mountains: Kemp and McCaig 1984; Jones and Robertson 1991; but no equivalent formation exists in the Othris Mountains), 4. Upper Cretaceous platform carbonates (Orliakas group in Northern Pindos Mountains: Jones and Robertson 1991); 5. Uppermost Cretaceous—Eocene flysch (Pindos Flysch in Northern Pindos Mountains: Richter 1976 (the latter two are not distinguished on the geological map of Othris by Ferrière (1982)).

Northern Pindos Mountains

Avdella *Mélange*

Traditionally, the Avdella *Mélange* is considered as a tectono-stratigraphic complex, which includes massive extrusive rocks (e. g. pillow basalts), various pelagic sediments and metamorphic blocks (Jones and Robertson 1991; Jones et al. 1992). This lower “volcanosedimentary” unit (e.g. ophiolitic olistostrome/*mélange*) was assigned a Triassic age using radiolarian biostratigraphic data (De Wever et al. 1979). Samples that yielded Triassic radiolarians were collected from Kandila Hill, northwest

Liagkouna Mountain, close to Armata and near to Alatopetra (Fig. 2). Poorly preserved radiolarian assemblages was assigned a Carnian–Norian age, based on the occurrence of *Capnuchosphaera* spp. and *Xiphothecaella* spp. This age is supported by a detailed revision of the family Capnuchosphaeridae (Kozur et al. 2009) and the family Xiphothecaellidae (Kozur et al. 2007). The *mélange* is overthrust by Jurassic ophiolitic complex (Fig. 2), which is unconformably overlain by Upper Cretaceous rudist limestone in some areas. Thrust sheets and blocks of sedimentary rock within the *mélange* include pelagic carbonates, cherty limestones and radiolarian cherts. The matrix of these blocks and dismembered thrust sheets is mainly shale, marl and sandstone (Jones and Robertson 1991).

Section NE of Avdella, radiolarite overlying pillow basalt

The basalt/radiolarite section is a block within the *mélange* with a sheared black shale matrix and is exposed on the northern side of a small stream gorge where it crosses the road (Fig. 3.1–5). Green pillow lava is overlain by a 6–7 m thick, steep to subvertical succession of red, ribbon radiolarite (Fig. 3.4). Eight samples (I/7/1–I/9/3) were collected from this section.

Section on the southeastern ridge of Mt. Straggopetra (including Mt. Lekani)

It includes green, amygdaloidal basalt lava ~300 m in thick with upper Scythian to lower Anisian-aged slide blocks as “olistoliths” (Migiros and Tselepidis 1990, Fig. 2) of Hallstatt-type pelagic limestone. Overlying red radiolarites, several tens of meters thick form the summit of Mt. Lekani (Fig. 5). Apparently, the contact between basalts and red radiolarites is continuous, but mostly covered by grass. A close intermixing between reddish to pinkish micrite and amygdaloidal pillow basalt of the same color provides evidence that the lime mud was unlithified during extrusion of the lava (Fig. 6.3–5). Hallstatt-type limestone blocks ranging in size from <10 cm in across tens of meters in size (to 10–12 m in height and 40–50 m in length) appear to have been emplaced together in one gravitational emplacement event. The overlying red radiolarite outcrop beneath the summit of Lekani Mt. and southeastwards toward the road connecting Lavda to Avdella. Three samples (Av-3,4,6) were collected from this section.

Section NW of Avdella village

The outcrop comprises large blocks with intercalations of slightly chloritized basalt, pillow basalt with red limestone, mudstone, chert and radiolarite (Fig. 6.1). Four samples

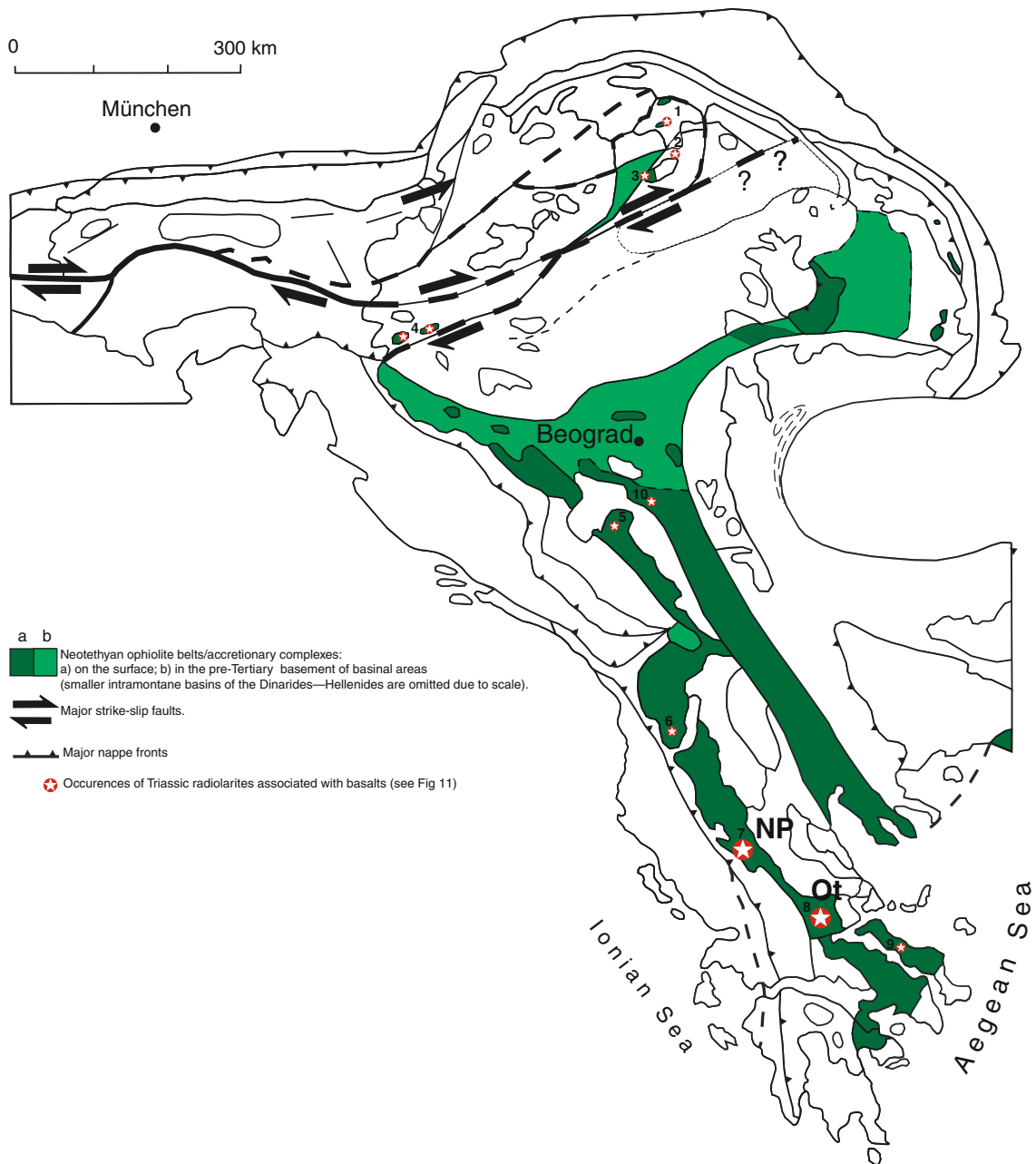


Fig. 1 Neotethyan ophiolite belts in the Alps, Carpathians, Dinarides and Hellenides and their dispersed fragments (= small-sized “*disrupted terranes*” of Neotethyan oceanic origin) in the Circum-

Pannonian region. (Base map simplified after Kovács et al. 2000). NP North Pindos, Ot Othris

(G.4, G.4/1-3) were taken from the limestone and radiolarite sequence.

Othris Mountains area

Loggitsion unit

In the central part of Othris Mountain area (Fig. 7), thrust sheets and blocks of pillow basalts are overlain by (or locally

interfinger with) ribbon radiolarites and cherts (Ferrière 1976, 1982; Bortolotti et al. 2008). The blocks usually occur in a strongly sheared black shale matrix. The radiolarites are dated as Middle to Upper Triassic (upper Anisian–lower Ladinian to Norian) in age (Bortolotti et al. 2008).

Section S of Metallion

In its lower part of the section dark green, basalt whit sheeted dykes is exposed (up to a few meters above the

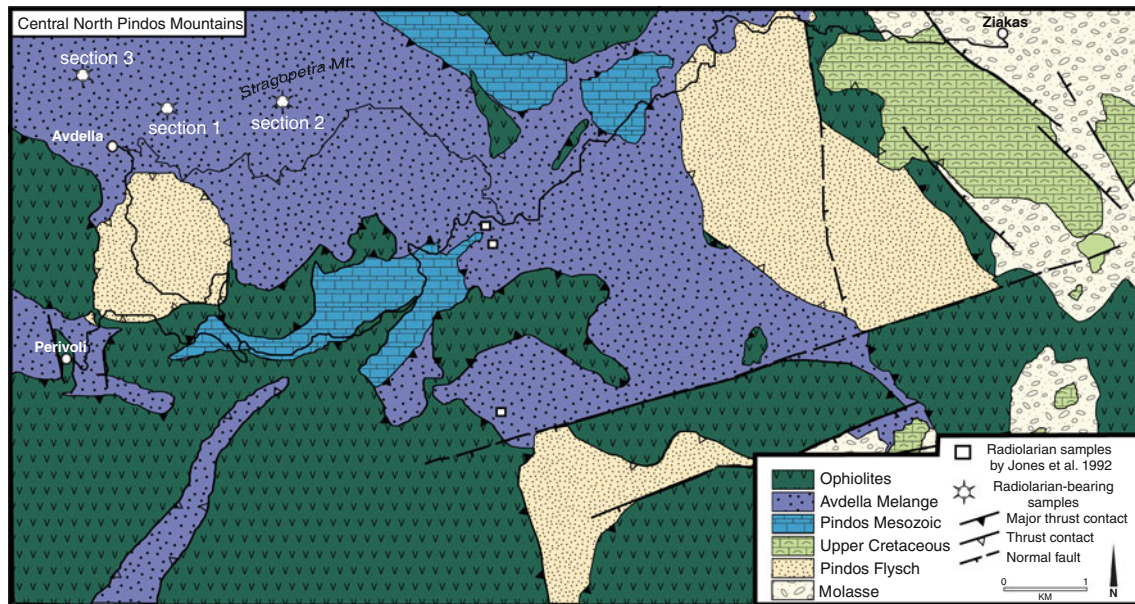


Fig. 2 Geological sketch map of the central part of Northern Pindos Mountains (area between Avdella–Perivoli–Ziakas villages, W of Grevena town), after Jones and Robertson (1991), slightly modified. Section 1 (N:40°00′44.30; E:21°08′02.48) = Section NE of Avdella,

radiolarite overlying pillow basalt; Sect. 2 (N:40°00′52.31; E:21°09′48.07) = Section on the southeastern ridge of Mt. Stragopetra (incl. Mt. Lekani); Sect. 3 (N:40°01′24.5; E:21°05′34.8) = Section NW of Avdella village

Fig. 3 Avdella Mélange, Northern Pindos Mountains. Gorge of a small creek NE of Avdella village, opening to the forestry road encircling the head of the Avdella Valley. 1: The section is exposed ~1.5 km to the NE of Avdella village (Fig. 2), along a forestry road paralleling the valley. Opening of the gorge to the forestry road, with the small creek in the middle, exposing the sheared black shale matrix of the mélangé in the back part. 2: Typical mélangé exposed in the back part of the gorge, with sheared black shale matrix and pinkish micritic limestone, greenish basalt blocks. 3: Lower part of the investigated red radiolarite section. 4: The investigated radiolarite section. 5: The upper part of the investigated section, with a fault cutting across. 6: Contact between the uppermost part of the underlying green basalt and the base of the overlying red radiolarite



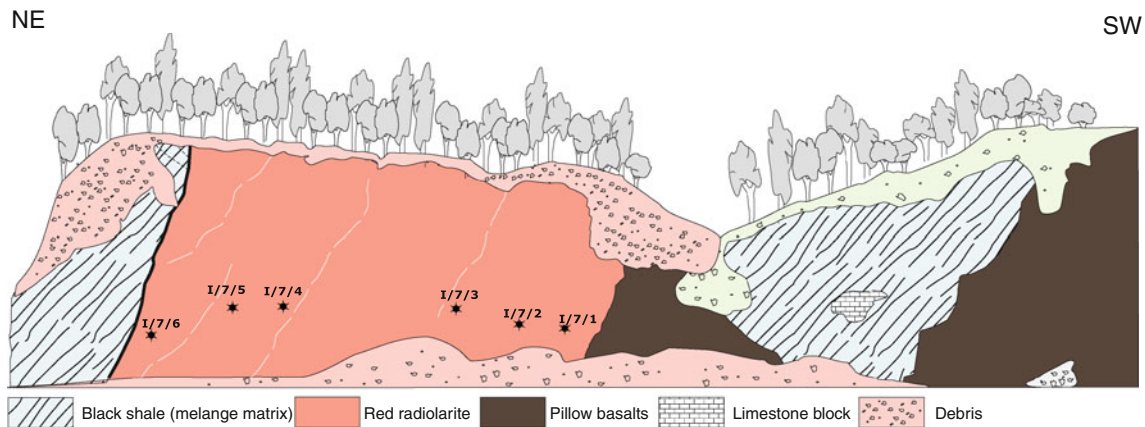
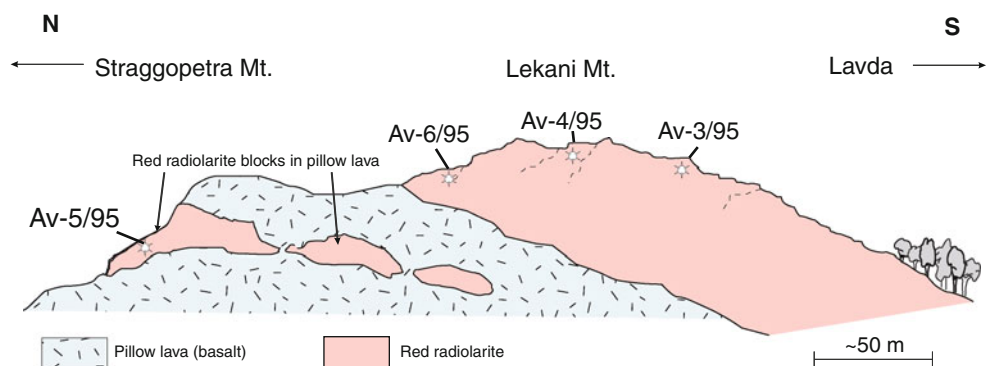


Fig. 4 Section NE of Avdella, exposing radiolarite overlying pillow basalts in Avdella Mélange, Northern Pindos Mountains

Fig. 5 This section is located on the southeastern ridge of Mt. Straggopetra (1,656 m), at an altitude of ~1,400 m above sea level



road), overlain by poorly exposed black shale (1–2 m) and red radiolarite (2–3 m) (Figs. 6.2; 8).

Radiolarian fauna and biostratigraphy

Radiolarians were extracted from the cherts and red radiolarites using hydrofluoric acid and standard laboratory techniques (De Wever et al. 2001). Preservation is generally poor, sometimes very poor, but a few identifiable taxa have been found suitable for biostratigraphic dating (Figs. 9, 10). Table 1 summarizes the taxa identified in samples from the sections.

Preliminary results of our investigations (Ladinian–Carnian age of radiolarites) are mentioned in Skourtsis-Coroneou et al. 1995. The first sample (I/7/1) of the section NE of Avdella, which was taken just a few centimeters above a basaltic block (Fig 3.6), yielded numerous unidentifiable spongy shells and a few slightly curved, apical polar spines of an Oertlispongidae species. Presumably belonging to the genus *Paroertlispongia* KOZUR and MOSTLER, 1981. Several genus of the family Oertlispongidae KOZUR and MOSTLER (in Dumitrică et al. 1980) were the stratigraphically most important radiolarian genus

in the Tethyan realm during the Middle Triassic (Kozur and Mostler 1996). The stratigraphic range of the genus *Paroertlispongia* extended from the Middle Anisian (Bithynian) to the Early Carnian. We have identified *Paroertlispongia multispinosus* KOZUR and MOSTLER in the *Tetraspinocyrtis leavis* radiolarian zone (correspond with *Paraceratites trinodosus* Ammonoid Subzone—lower Illyrian, Anisian) in the Felsőörs section of the Balaton Highland in Hungary (Ozsvárt and Dosztály 2007). The recently accepted GSSP (Global Boundary Stratotype Section and Point) of the Anisian–Ladinian stage boundary (Brack et al. 2003, 2005) places the base of the Ladinian at the base of the *Curioni* Ammonoid Zone. Consequently, based on the presence of the species *Paroertlispongia multispinosus* KOZUR and MOSTLER, this sample could be assigned to the upper Anisian (lower Illyrian) to the lower Ladinian (upper Fassanian). In addition, based on the presence of *Eptingium manfredi* DUMITRICĂ, sample I/7/2 can be assigned to the upper Anisian (lower Illyrian) to lowermost Ladinian (lower Fassanian). However, the presence of *Muelleritortis firmum* (GORIČAN) introduces some ambiguity in the age assignment. Originally, *M. firmum* (GORIČAN) was reported from the Longobardian of the Vršič and Mokronog sections in Slovenia (Goričan and

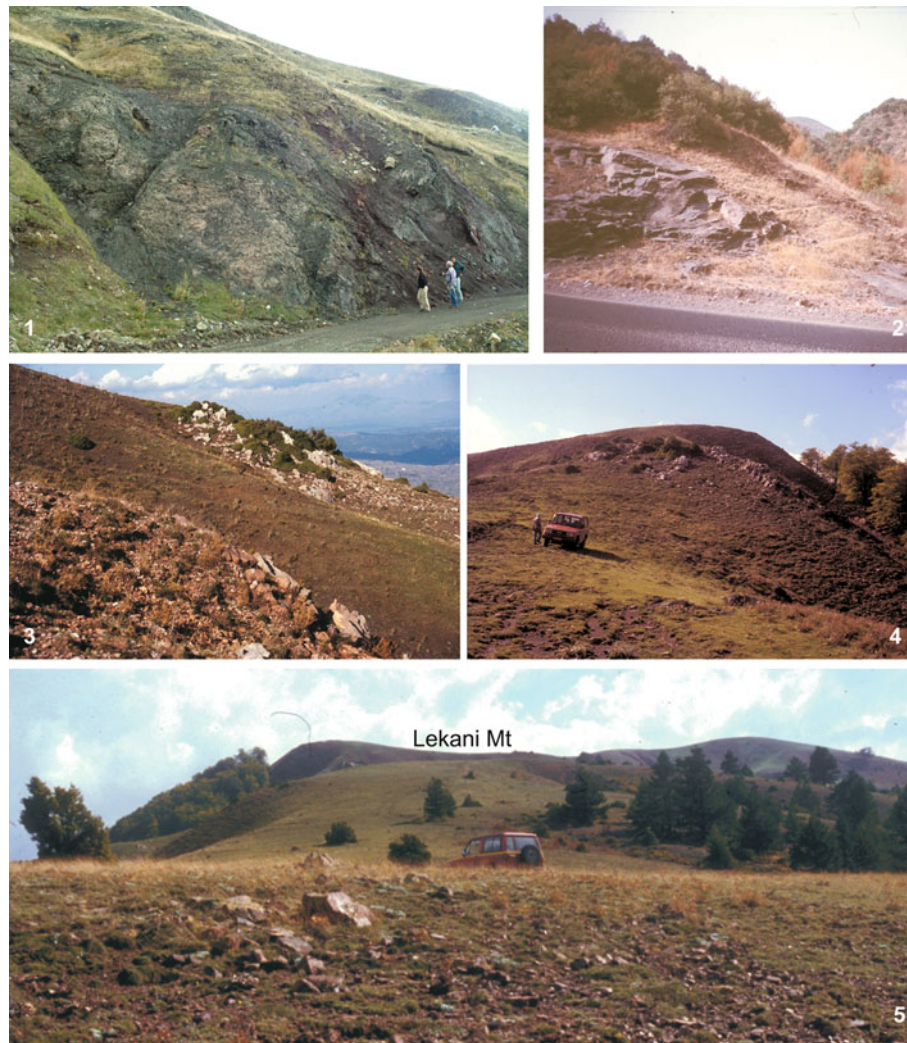


Fig. 6 Avdella Mélange, Northern Pindos Mts. and Loggitsion Unit, western part of Othris Mts. *1*: This section is located 3.5 km northwest of Avdella village. *Dark, purplish red* radiolarite (locality of sample G.4) and *red*, micritic limestone wedged between a *reddish* and *greenish*, amygdaloidal basalt block (*left*) and a serpentinite block (*right*). The latter also includes red micritic limestone fragments, tectonically reworked along a shear zones. NNE of Avdella village, along the forestry road encircling the head of the Avdella valley. *2*: Loggitsion Unit, western part of Othris Mts.: *Green* basalt, with some black shale, than *red* radiolarite (*behind the bush*) above it, yielding

Ladinian radiolarians (described in Ozsvárt et al. 2006). Along the western side of the road leading from Domokos to Lamia, 5 km S of Metallion village. *3–4*: Slide blocks of reddish, micritic Hallstatt-type limestone in amygdaloidal basalt “matrix” on the southeastern extension of Straggopetra Mt. From the *lower* part of the limestone block shown on Fig. 4, Migiros and Tselepidis (1990) published an *Upper Scythian* ammonoid fauna. *5*: *Red* radiolarite succession of Lekani Mt. overlying the amygdaloidal basalt sequence of the SE extension of Straggopetra Mt

Buser 1990). This species co-occurs with *Triassocampe scalaris* DUMITRICĂ, KOZUR and MOSTLER, which is known from the upper Anisian (lower Illyrian) to the lowermost Ladinian (lower Fassanian) beds within the entire Eurasian Tethyan realm (Kozur and Mostler 1994). The most probable stratigraphic age range of *M. firmum* (GORIČAN) is therefore upper Anisian (lower Illyrian) to lowermost Ladinian (lower Fassanian), rather than Longobardian. In summary, sample I/7/2 is regarded as being upper Anisian (lower Illyrian) to lowermost Ladinian (Lower Fassanian).

Sample I/7/3 yielded poorly preserved radiolarians. The presence of the species *Annulotriassocampe* cf. *sulovens* KOZUR and MOCK and *Capnuchosphaera* sp. suggests a younger age than lower Carnian for this bed, because the range of the genus *Annulotriassocampe* is Middle Anisian–Middle Norian, while the genus *Capnuchosphaera* first appears in the early Carnian in the Tethyan realm and disappears in the Middle Norian.

Samples I/7/4–6 were collected a few meters above basalt. The poorly preserved radiolarian assemblage of this samples

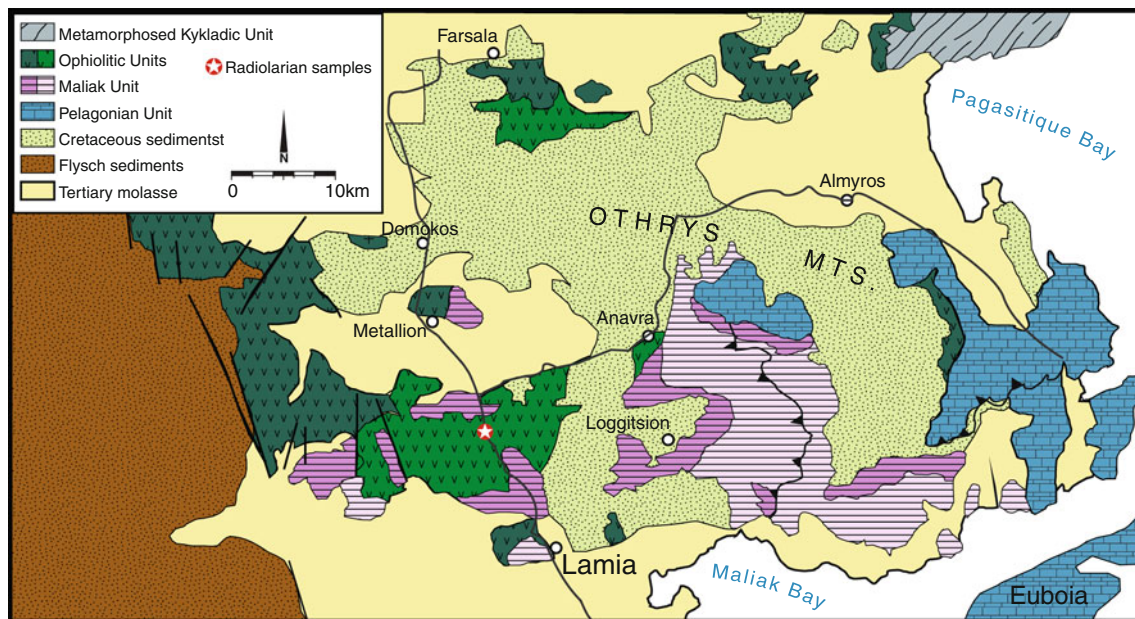


Fig. 7 Geological sketch map of the Othrys Mts. (modified after Ferrière 1982), with location of the outcrop studied with radiolarite above basalt, shown on Fig. 3.2 (along the road between Metallion and Lamia)

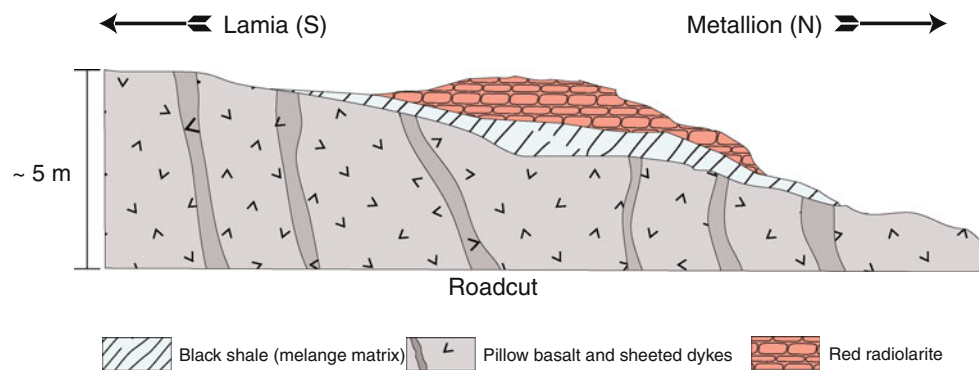


Fig. 8 The outcrop is located along the western side of a road leading from Domokos to Lamia, 10 km S of Metallion (Fig. 7), near the boundary with the dykes of the Loggitsion and ophiolite units

suggests a Norian age on the basis of the co-occurrence of *Capnuhosphaera crassa* YEH, which first appears in the Carnian, and of *Capnodoce anapetes* DE WEVER which is known from the lower Norian beds in the Antalya Nappes (De Wever et al. 1979; De Wever 1982).

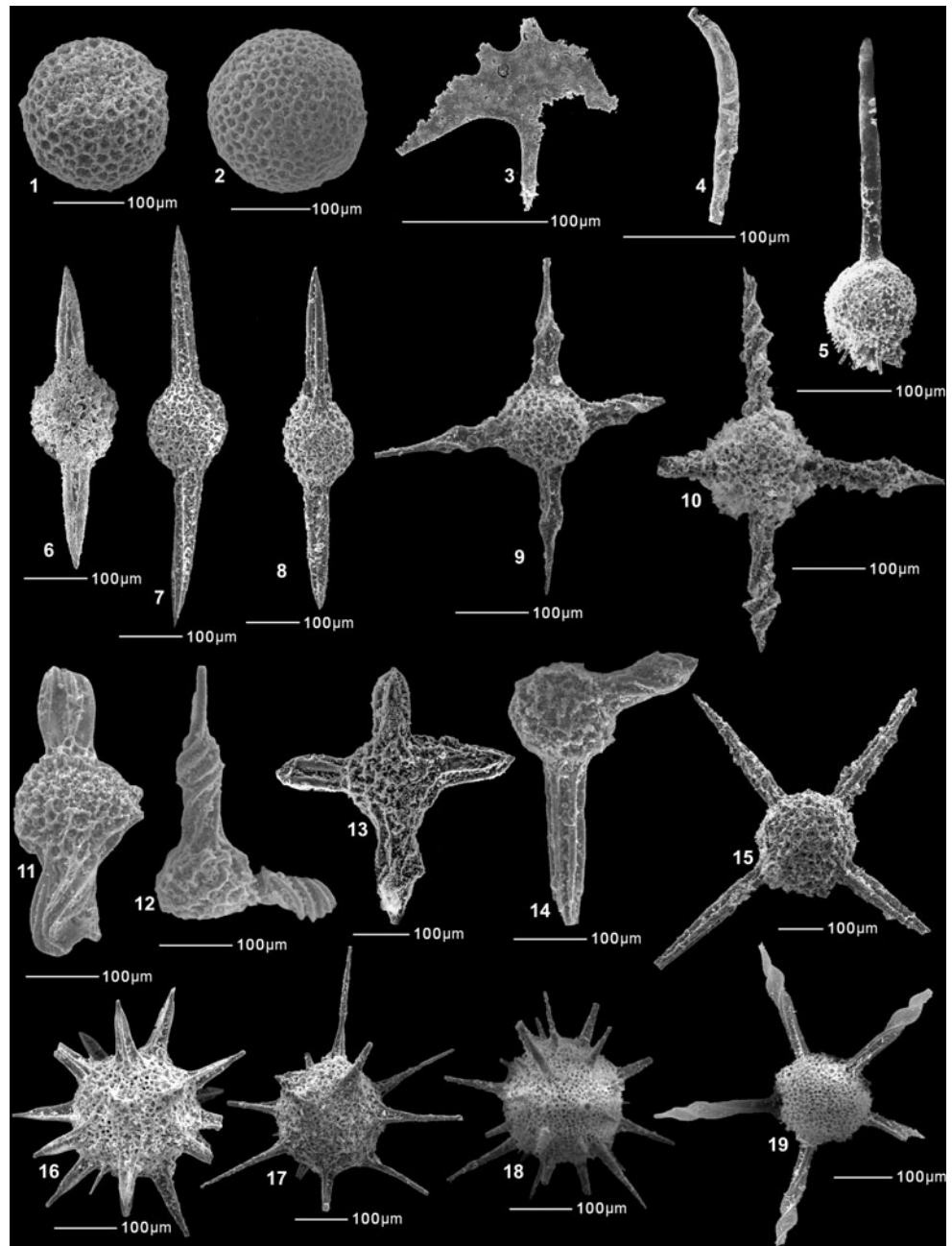
In summary, the lower part (sample I/7/1-2) of the red radiolarite succession exposed 1.5 km to the NE of Avdella village is certainly of Upper Anisian (Lower Illyrian) to lowermost Ladinian (Lower Fassinian) age. The middle part of the succession (Sample I/7/3) is probably Upper Ladinian, and the upper part (Sample I/7/4-6) Upper Triassic (probably Carnian to Lower Norian).

Poorly preserved radiolarian assemblages were also extracted from the three samples (Av-6/95, Av-4/95, Av-3/95) from the section on the southeastern ridge of Mt. Straggopetra (Table I). The probable age for the lower part

of this section (Av-6/95) is Ladinian (see above for a discussion of *Oertlispongidae*). Based on the presence of some poorly preserved *Capnuhosphaeridae*, an Upper Triassic age (probably Carnian to Norian) is assumed for the upper part of the section (Av-4/95, Av-3/95), because the common occurrence of the genus *Capnuhosphaera* is characteristic of the Carnian (Tuvalian) and is also very frequent in the Norian. Our data, in agreement with Migiros and Tselepidis (1990), show that the age of the amygdaloidal pillow basalt is between upper Scythian (base of Ammonite fauna; Migiros and Tselepidis 1990) and upper Anisian (or lowermost Ladinian).

Samples G.4, G.4/1–G.4/3 were collected from the section northwest of Avdella village where altered basalt is intercalated with red limestone, mudstone-chert and radiolarite. The co-occurrence of *Astrocentrus latispinosus*

Fig. 9 Illustration of selected radiolarian taxa from the Pindos Mts. 1—*Cenosphaera* sp. 1, G.4; 2—*Cenosphaera* sp. 2, G.4; 3—*Steigerispongus* sp., G.4/1; 4—*Paroertlispongus* cf. *weddegi* (Lahm), I/7/1; 5—*Paroertlispongus multispinosus* Kozur and Mostler, I/7/2; 6—*Pseudostylosphaera coccostyla acrior* (Bragin), I/7/2; 7—*Pseudostylosphaera longispinosa* Kozur and Mostler, G.4; 8—*Pseudostylosphaera longobardica* Kozur and Mostler, G.4/1; 9—*Plafkerium? muelleri* Dumitrică, Kozur and Mostler, G.4; 10—*Plafkerium? nazarovi* Kozur and Mostler, I/7/2; 11—*Muelleritortis* sp. 1, G.4; 12—*Muelleritortis* sp. 2, G.4; 13—*Muelleritortis firmum* (Goričan and Buser), I/7/2; 14—*Tritortis* sp. G.4/2; 15—*Parasepsagon asymmetricus* Kozur and Mostler, I/7/1; 16—*Astrocentrus latispinosus* (Kozur and Mostler), G.4/3; 17—*Astrocentrus pulcher* Kozur and Mostler, G.4/1; 18—*Triassospongosphaera multispinosa* (Kozur and Mostler), G.4/3; 19—*Pentaspogodiscus ladinicus* Dumitrică, Kozur and Mostler, G.4/1



(KOZUR and MOSTLER), *Holzmadia reticulata* DUMITRICĂ, KOZUR and MOSTLER, *Plafkerium muelleri* DUMITRICĂ, KOZUR and MOSTLER, *Stauracontium trispinosum* (KOZUR and MOSTLER), *Pseudostylosphaera longispinosa* KOZUR and MOSTLER (see Table 1) suggests a Middle Triassic age (Illyrian–Longobardian) for the lower part of the section. The taxa identified from samples G.4/1–G.4/3: *Astrocentrus pulcher* KOZUR and MOSTLER, *Pseudostylosphaera longobardica* KOZUR and MOSTLER, *Steigerispongus* sp., *Triassospongosphaera multispinosa* (KOZUR and MOSTLER), see Table 1, suggest a Longobardian age.

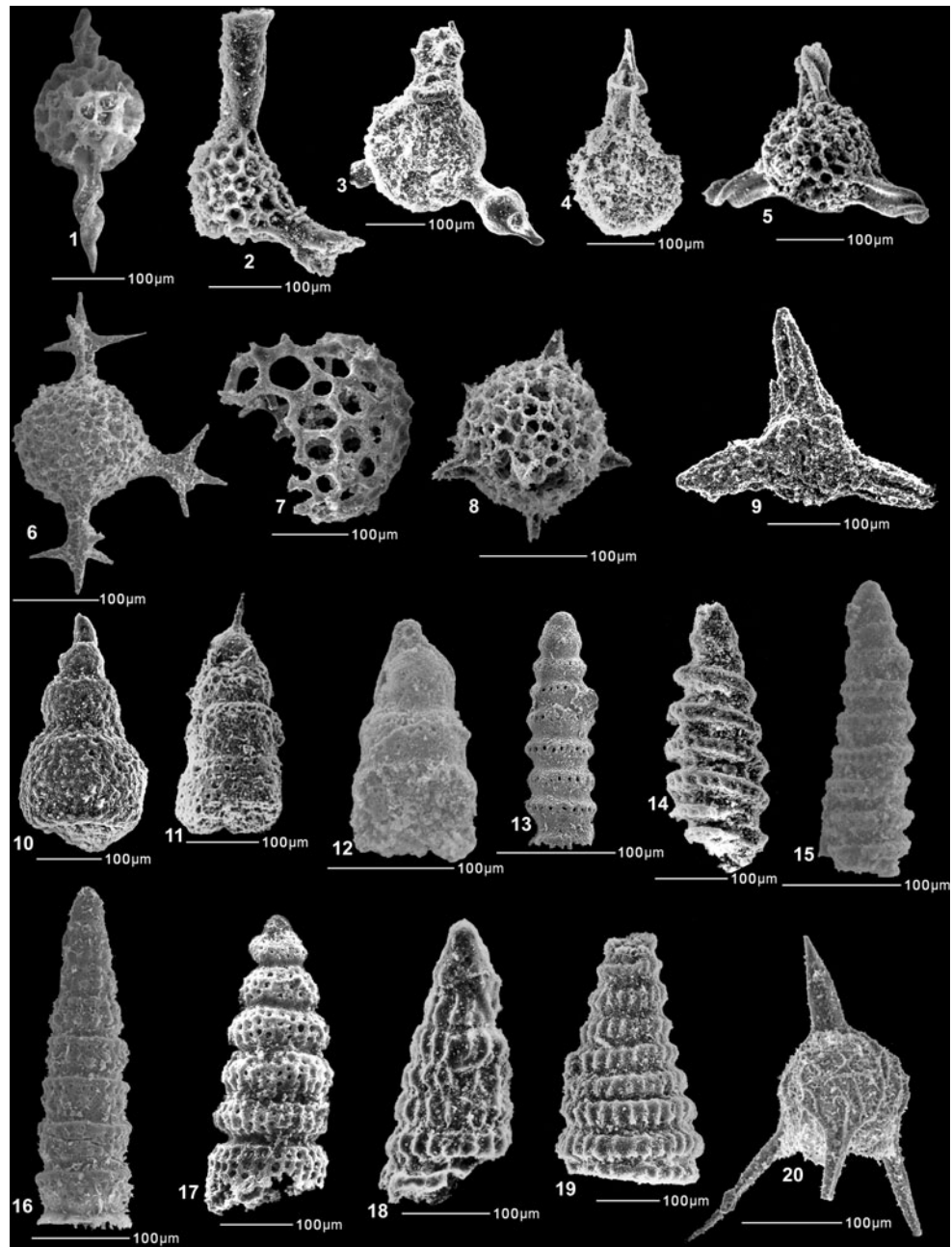
Sample II/3/1 was collected in the Othris Mountain section (Fig. 8), and a Middle Triassic (probably Illyrian) age is proposed based on the presence of moderately preserved radiolarians (see Table 1).

Discussion

Using radiolarians, we have dated four radiolarite outcrops within tectono-stratigraphic mélanges of the Northern Pindos and Othris Mountains beneath the Jurassic

Fig. 10 Illustration of selected radiolarian taxa from the Pindos and Othris Mts.

1—*Pantanelium* sp., Av-4;
 2—*Capnodoce* cf. *anapetes* De Wever, I/7/5; 3—*Capnuchosphaera* sp. 1, II/4/5;
 4—*Capnuchosphaera* sp. 2, II/4/5; 5—*Sarla* sp., II/7/3;
 6—*Stauracantium trispinosum* (Kozur and Mostler), G.4/1;
 7—? *Pentactinocarpus* sp., G.4/1; 8—*Welirella* sp., G.4/3;
 9—*Eptingium manfredi* Dumitrică, I/7/2;
 10—? *Nakasekoellus* sp., Av-3;
 11—*Anisicyrtis* sp., I/7/1;
 12—*Anisicyrtis* sp., I/7/1;
 13—*Annulotriassocampe campanilis* Kozur and Mostler, I/7/1; 14—*Annulotriassocampe* cf. *sulovensis* (Kozur and Mock), I/7/3 slightly deformed;
 15—*Annulotriassocampe* cf. *sulovensis* (Kozur and Mock), I/7/3; 16—*Triassocampe* cf. *scalaris* Dumitrică, Kozur and Mostler, II/3/1;
 17—*Pararuesticyrtium* sp., I/7/3; 18—*Corum kraineri* Tekin, I/7/6; 19—*Whalenella* sp., II/4/5; 20—*Holzmadia reticulata* Dumitrică, Kozur and Mostler, G.4/1



ophiolites in Northern Greece, two of them directly overlying basalts. The latter two near Avdella village in Northern Pindos Mts. provide evidence that begin already before the earliest Ladinian time. In particular, the Stragopetra—Lekani Mt. section shows that an ~300 m thick amygdaloidal basalt succession had already erupted during the Anisian. The geochemical study on a single sample from this section shows that the within-plate (WP) origin is the most probable for these basaltic rocks (Kiss et al.). In addition, detailed geochemically analysis indicated the WP and MOR-type basalts from several localities in Avdella Mélange (Jones and Robertson 1991), and WPB and

N-MORB from Othris Mountains (Bortolotti et al. 2008). Field evidence suggests that basaltic volcanism started on deep water environment (e. g. basalts intensively mixed with pelagic lime mud resulting in a characteristic peperitic facies) which might concerned to the initial stage of rifting of the Neotethys in the Hellenide sector or might connected to a seamount on the oceanic floor, although we do not know any characteristic slope or terrigenous sediments in this area, temporarily. The latter geodynamic model seems unlikely, as it is not probable, that in the initial stage of rifting there was enough space away from the rift zone for larger seamounts to have formed. Upper Scythian

Table 1 Middle and Upper Triassic radiolarian occurrences in Northern Pindos Mts. and in Othrys Mts

	Northern Pindos Mountains						Othrys Mountains										
	Section NE of Avdella			Straggopetra Mt.			Section NW of Avdella village										
	Illyrian-Fassanian	?Carnian	?Carnian-Norian	Anisian-Ladinian	?Carnian	Longobardian	Roadcut (S of Metallion)										
I/7/1	I/7/2	I/7/3	I/7/4	I/7/5	I/7/6	I/8/3	I/9/3	Av-4	Av-3	G.4	G.4/1	G.4/2	G.4/3	II/3/1	II/4/5	II/7/3	
<i>Anisicyrtis</i> sp.	+																
<i>Annulotriassocampe campanilis</i> Kozur and Mostler	+																
<i>Annulotriassocampe cf. sulovens</i> (Kozur and Mock)			+														
<i>Archaeospongoprunum mesotriassicum</i> Kozur and Mostler																	
<i>Astrocentrus latispinosus</i> (Kozur and Mostler)																	
<i>Astrocentrus pulcher</i> Kozur and Mostler																	
<i>Capnodoce cf. anapetes</i> De Wever																	
<i>Capnodoce</i> sp.																	
<i>Capnucho-sphaera crassa</i> Yeh																	
<i>Capnucho-sphaera</i> sp. 1																	
<i>Capnucho-sphaera</i> sp. 2																	
<i>Cenosphaera</i> sp. 1																	
<i>Cenosphaera</i> sp. 2																	
<i>Corum kraineri</i> Tekin																	
? <i>Corum</i> sp.																	
<i>Cryptostephandium</i> sp.																	
<i>Eptingium manfredi</i> Dumitrićă																	
<i>Hozmadia reticulata</i> Dumitrićă, Kozur and Mostler																	
<i>Hozmadia</i> sp.																	
? <i>Monocapnucho-sphaera</i> sp.																	
<i>Muelleritortis firmum</i> (Goričan and Buser)																	
<i>Muelleritortis</i> sp. 1																	
<i>Muelleritortis</i> sp. 2																	
? <i>Nakasekoellus</i> sp.																	
<i>Oertlispongos</i> sp. 1																	
<i>Oertlispongos</i> sp. 2																	
<i>Pantanelium</i> sp.																	
<i>Pararuesticyrtium</i> sp.																	
<i>Parasepsagon asymmetricus</i> Kozur and Mostler																	
<i>Paroertlispongos cf. weddigei</i> (Lahm)																	

Table 1 continued

	Northern Pindos Mountains						Othrys Mountains						
	Section NE of Avdella			Stragopetra Mt.			Section NW of Avdella village		Roadcut (S of Metallion)				
	Illyrian-Fassanian	?Carnian	?Carnian-Norian	Av-6	Anisian-Ladinian	?Carnian	G.4	G.4/1	G.4/2	G.4/3	II/3/1	II/4/5	II/7/3
<i>Paroerlispongius multispinosus</i> Kozur and Mostler	+												
<i>Paroerlispongius</i> sp.	+												
? <i>Pentactinocarpus</i> sp.													
<i>Pentaspiongodiscus ladinicus</i> Dumitriță, Kozur and Mostler													
<i>Plafkerium ? muelleri</i> Dumitriță, Kozur and Mostler													
<i>Plafkerium ? nazarovi</i> Kozur and Mostler	+												
<i>Pseudostylosphaera coccostyla acrior</i> (Bragin)	+												
<i>Pseudostylosphaera longispinosa</i> Kozur and Mostler													
<i>Pseudostylosphaera longobardica</i> Kozur and Mostler													
<i>Pseudostylosphaera</i> sp.	+												
<i>Sarla</i> sp.													
<i>Spongopallium</i> sp.													
<i>Stauracantium trispinosum</i> (Kozur and Mostler)													
<i>Steigerispongius</i> sp.													
<i>Stichocapsa</i> cf. <i>robusta</i> Matsuoka													
<i>Tiborella magnidentata</i> Dumitriță, Kozur and Mostler													
<i>Triassocampe</i> cf. <i>scalaris</i> Dumitriță, Kozur and Mostler													
<i>Triassocampe</i> sp.	+												
<i>Triassocamposphaera multispinosa</i> (Kozur and Mostler)													
<i>Tritortis</i> sp. 1													
<i>Tritortis</i> sp. 2													
<i>Welirella</i> sp.													
<i>Whalenella</i> sp.													

ammonoids and lowermost Anisian conodonts from a Hallstatt-type limestone block in the basal part of this Straggopetra—Lekani Mt. section (Migiros and Tselepidis 1990) suggest a lower age limit of the characteristic peperitic basaltic volcanism (Kovács et al. in press).

Having compared other areas, upper Illyrian (Anisian) to upper Ladinian radiolarians have been reported (partly from outcrops, where radiolarites are associated with basalts) from several localities in SE Slovakia (Meliaticum *sensu stricto*, Jaklovce subunit) (Mello et al. 1995) and the Tornakápolna unit of NE part of Hungary (Kozur and Réti 1986; Dosztály and Józsa 1992, Fig. 11). Upper Illyrian to lower Carnian radiolarian assemblages are also recorded in

the Darnó and Mónosbél units (De Wever 1984; Dosztály and Józsa 1992) of NE Hungary. In addition, upper Ladinian to Middle Norian radiolarian assemblages were dated from Zagorje region (NW Croatia) where radiolarites directly overlie basaltic volcanics (Halamić and Goričan 1995; Halamić et al. 2001; Goričan et al. 2005) (Fig. 11). Within the Dinaric Ophiolite Belt (DOB) of Serbia (for references see Hein and Obradović 1988; Kovács et al. in press), several localities (Zlatibor and Zlatar Mts.) include Ladinian to lower Norian radiolarites associated with basalts (Obradović and Goričan 1988; Vishnevskaya et al. 2009). Further localities were reported from the Vardar Zone Western Belt (Ovčar-Kabljar gorge, west of Čačak)

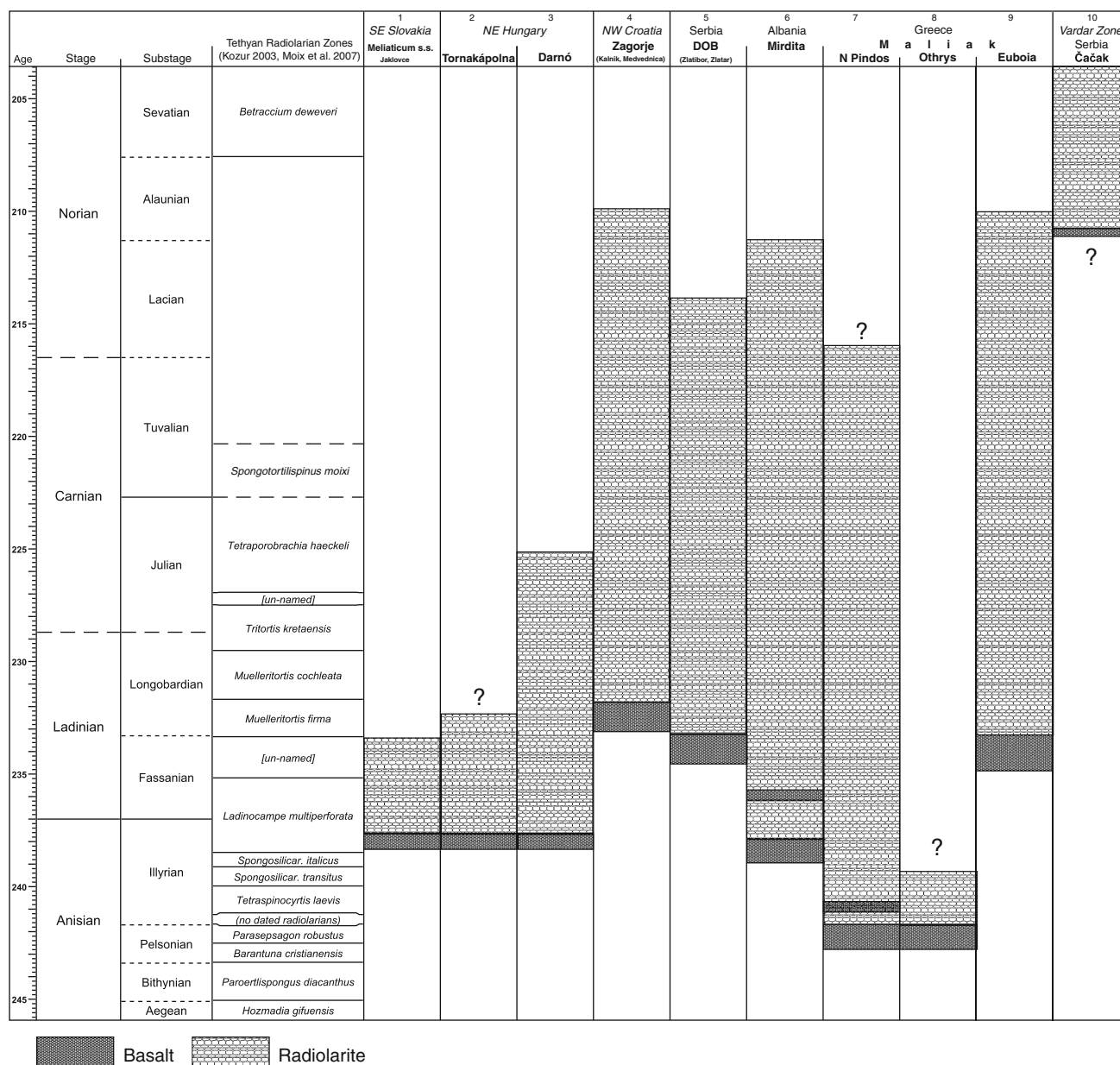


Fig. 11 Published age ranges of radiolarites with associated pillow basalts in the Western part of Mesozoic Neotethys. See text for References

and from DOB (Katusnica creek, near Gostilje, Zlatibor Mountains) where radiolarian-bearing rocks are associated with basalts, dated as Ladinian to Norian in age (Obradović and Goričan 1988; Vishnevskaya et al. 2009) (Fig. 11). Farther south, several localities in the Mirdita Zone in Albania have been shown to be of Anisian age (Fig. 11). Alkaline basalts within the mélange are directly overlain by upper Anisian (Illyrian) radiolarites (Chiari et al. 1994, 1996; Kellici et al. 1994; Marcucci et al. 1994; Bortolotti et al. 2006; Gawlick et al. 2008).

Conodont data from other parts of the Inner Hellenides are also relevant. Kaufmann (1976) has reported Upper Scythian red, Hallstatt-type limestones associated with basalts from Euboia Island. Upper Ladinian to ?Middle Norian radiolarian age data (Fig. 11) have also been reported from the Pagondas Mélange of Euboia Island (Danelian and Robertson 2001). As concerns conodonts, here we can add our own data from the Avdella Mélange: at the road branching W of Ziakas toward Avdella, we found Pelsonian (Middle Anisian) conodonts (*Gondolella hanbulogi*) in a ca. 2-m long, lens-like red limestone block within amygdaloidal basalt.

In the former area pelagic sedimentation began in the Late Scythian followed by basalt volcanism in the Late Scythian?—Anisian time. On the other hand, in the latter region, where remnants of the northwestern termination of the Neotethys ocean have been preserved, pelagic sedimentation began only in the Middle–Late Anisian time, followed by sporadically recorded basalt volcanism only in the Ladinian (Kovács 1984; Réti 1985; Lein 1987; Kozur 1991; Haas et al. 1995; Mello et al. 1995). Toward the northwest end of the Neotethys ocean, pelagic sedimentation began in the Middle–Late Anisian, followed by new oceanic crust forming in the Ladinian. Moreover, pelagic sedimentation (Hallstatt-type limestones, radiolarites) associated with basalt volcanism commenced in Oman already in the Middle–Late Permian (e. g. Maury et al. 2003). These age data clearly demonstrate the westwards propagation of Neotethyan rifting.

Conclusions

1. Radiolarites overlaying or intercalating with basalts from Northern Pindos Mountains and Othris Mountains, Inner Hellenides are dated as upper Anisian to lower Norian age.
2. The oldest radiolarians indicate Middle Triassic age (upper Anisian, probably lower Illyrian), which has been documented for the first time in Northern Pindos Mountains. This is the oldest radiolarian age data from this area.

3. The new data presented here constrain the beginning of basaltic volcanism in Western Greek Ophiolites Belt as pre-Ladinian (Late Scythian?–Anisian).
4. These biostratigraphic constraints suggest that Neotethyan rifting started earlier in the Hellenidic domain, than in the Eastern Alps–West Carpathian region, but significantly later than in Oman. This indicates a considerable time shift of the onset of Neotethyan rifting from east toward west.

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